

EMC-2015-001

Title: Multiscale investigation of perennial flow and thermal influence of headwater streams into fish bearing systems

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1. Introduction

The impacts of timber harvesting and other land uses on water quality have been an environmental concern for many years. One of the primary concerns is the potential for land use activities to produce increases in stream temperature during the summer and increases in suspended sediment yields during the wet season (Macdonald *et al.*, 2003; Gomi *et al.*, 2005; Moore *et al.*, 2005; Croke and Hairsine, 2006; Gomi *et al.*, 2006; Gravelle and Link, 2007). Water temperature and sediment are the two primary water quality constituents that have been recognized as the dominant causes of impairment in streams in Northern California and throughout the Pacific Northwest (ODEQ, 2004; Hanak *et al.*, 2011). Temperature, in particular, is of special concern in fish bearing streams, especially where threatened and/or endangered aquatic fish species are present. In California, the Forest Practice Rules (FPRs) afford the most protection to Class I (fish bearing) relative to Class II (aquatic life other than fish) and Class III streams (not supporting aquatic life). However, it has been recognized that headwater systems can be critically important to the water quality in downstream sites (MacDonald and Coe, 2007). This led to the establishment of provisions for Class II Large (Class II-L) watercourses according to the “Anadromous Salmonid Protection Rules, 2009”, and modified by the “Class II-L Identification and Protection Amendments, 2013” rule package approved by the State Board of Forestry and Fire Protection in October, 2013. One of the objectives of these rules is to protect anadromous salmonid habitat by minimizing potential increases in temperature, sediment, and nutrients from Class II and Class III watercourses draining into Class I systems. This amendment included an improved method to identify Class II-L watercourses based on drainage area and active channel width. A Class II-L watercourse is defined as a Class II watercourse that has either of the following characteristics:

- contributing drainage area of ≥ 100 acres in the Coast Forest District, or ≥ 150 acres for the Northern and Southern Forest Districts
- an average active channel width of 5 feet or greater near the confluence with the receiving Class I watercourse.

Given that the above method for determination will sunset on January 1, 2019 pending further evaluation (as specified in the amended FPRs), there is an urgent need to assess the efficacy of the rule.

Aside from the practical challenges associated with defining Class II-L systems, there are also scientific questions related to both the variability in the geographic extent of Class II-L systems and the assessment of their actual impacts on Class I systems. Indeed, the potential influence that Class II-L watercourses can impose downstream will likely depend on the hydrologic regime of the system, which is a function of physiographic and climatic variables, as well as its hydrologic connection to downstream reaches. Therefore, in this proposal we assume

that for Class II-L systems to potentially impact downstream reaches they must be perennial (i.e., hydrologically connected) during the time period of the potential negative effect. In the case of stream temperature, streams must be hydrologically connected during the summer period (i.e., perennial). Therefore, understanding the spatial and temporal variability of the perennial extent of streams is central to the rationale of this proposal.

The upper extent of perennial flow generally varies across the landscape as a function of natural landscape characteristics, climatic regimes, and land use impacts (Montgomery and Dietrich, 1989; Prosser and Abernethy, 1996; Tucker and Bras, 1998; Jaeger *et al.*, 2007; Costigan *et al.*, 2016). For instance, lithology appears to control source area size of forested streams in Washington underlined by basalt and sandstone (Jaeger *et al.*, 2007). However, this study also reported that in both lithologies the location of the channel heads and the start of perennial flow coincided. Channel heads can be defined morphologically as “the upslope limit of erosion and concentration of flow within steepened banks” (Montgomery and Dietrich, 1989). While field campaigns to map channel heads are time consuming and expensive, spatial analysis including digital elevation terrain derived methods offer a potential alternative that is less time consuming and more cost effective (Fritz *et al.*, 2008). The early DEM methods were based on a contributing area threshold or on a slope-area scaling relationship (Montgomery and Dietrich, 1989; Tarboton *et al.*, 1991). Recently, LiDAR information has been used in more refined and accurate schemes (Passalacqua *et al.*, 2010; Clubb *et al.*, 2014). However, not all regions have LiDAR data available yet. Therefore, comparative studies of the effects of improved DEM resolution over a range of environmental conditions could improve understanding of the limits of the different methods at identifying locations of channel heads and perennial flow. Initial research indicates that the discrepancy between the outcomes from different data resolution increases with terrain steepness (Grieve *et al.*, 2016). Additionally, it is critical to characterize the hydrologic regime (e.g., perennial vs. ephemeral) of a stream, as this can influence the accurate identification of channel heads and, therefore, the efficiency of terrain analysis (Wang and Wu, 2013; Costigan *et al.*, 2016).

From the perspective of this investigation, the perennial extent of a stream is a primary control on the transmission of thermal inputs downstream as cumulative effects (Beschta and Taylor, 1988; Gregory *et al.*, 1991). If a headwater stream warms, this impact is only relevant to downstream reaches as long as they are hydrologically connected. Given the importance of stream temperature as a water quality parameter, there have been many studies regarding changes in the thermal regimes of streams following forest management activities (Brown, 1969; Brown and Krygier, 1970; Macdonald *et al.*, 2003; Moore *et al.*, 2005; Dent *et al.*, 2008; Kibler *et al.*, 2013; Guenther *et al.*, 2014; Bladon *et al.*, 2016). Additionally, there is also concern about the downstream transmission of heated water, which would increase the spatial extent of thermal effects on aquatic ecosystems (Moore *et al.*, 2005). This concern has been reinforced by observations of heat inputs being transmitted downstream as cumulative effects (Beschta and Taylor, 1988; Gregory *et al.*, 1991). As such, asymptotic warming is often the supported conceptual paradigm for longitudinal stream temperature patterns (Caissie, 2006). However, the general model of downstream warming likely oversimplifies stream temperature dynamics (Dent *et al.*, 2008; Leach and Moore, 2011). Studies from California, Oregon, British Columbia, and elsewhere have also demonstrated both natural stream cooling in a downstream direction (Madej *et al.*, 2006; Fullerton *et al.*, 2015), as well as cooling of warmed water flowing out of a clearcut

and back into a closed-canopy section (McGurk, 1989; Keith *et al.*, 1998; Story *et al.*, 2003). In the context of timber harvesting, recent research from across 10 sites in Oregon showed that lithology can be a dominant control on the downstream effect in terms of not only sediment (Bywater-Reyes *et al.*) but also temperature (unpublished data). However, while there is growing appreciation for the high degree of variability in longitudinal stream temperature dynamics (Ebersole *et al.*, 2003; Davis *et al.*, 2015; Fullerton *et al.*, 2015; Louen, 2016), it is increasingly important to evaluate the influence of local and regional drivers on downstream thermal regimes to improve our ability to predict disturbance responses. This is particularly critical for streams in California given its geologic and geomorphological complexity, as well as its warmer climate regime compared to that in the PNW.

2. Objectives

This proposed research is consistent with the monitoring priorities outlined in the California Board of Forestry and Fire Protection's Effectiveness Monitoring Committee's (EMC) Strategic Plan¹. The EMC's Strategic Plan identifies Class II-L monitoring as a priority for evaluating the effectiveness of the California Forest Practice Rules in protecting, maintaining, and restoring riparian function in larger non-fish bearing watercourses. Specifically, the objectives of this proposal are:

- a) Investigate the variability of the relationship between drainage area, active channel width, and perennial flow extent across the Anadromous Salmonid Protection (ASP) area (Fig. 1);
- b) Compare the relationships derived in (a) to the rule criteria for Class II-L identification in terms of both drainage area and average active channel width; determine if these criteria are effective in identifying perennial Class II-L watercourses in different lithologies, or if rule modifications are needed;
- c) Conduct a pilot study to investigate the downstream propagation of water temperature from Class II-L systems in sites with contrasting lithology

To achieve the objectives we propose to integrate spatial analysis, field observations and mapping, and coupled measurements of stream and air temperature in a two level multiscale approach at the regional and catchment scales. We will utilize available geospatial information including digital elevation terrain models, LiDAR, and geology and hydrometric data to stratify a regional field campaign representative of the ASP region (Fig. 1). In addition, we plan a process driven investigation of the longitudinal variability of the relation between air and stream temperature considering 2 sites representative of the geologic variability of the region.

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http://bofdata.fire.ca.gov/board_business/binder_materials/2017/jan_2017/full/full_12.0_b__2__emc_strategic_plan_clean_version_01_19_17.pdf

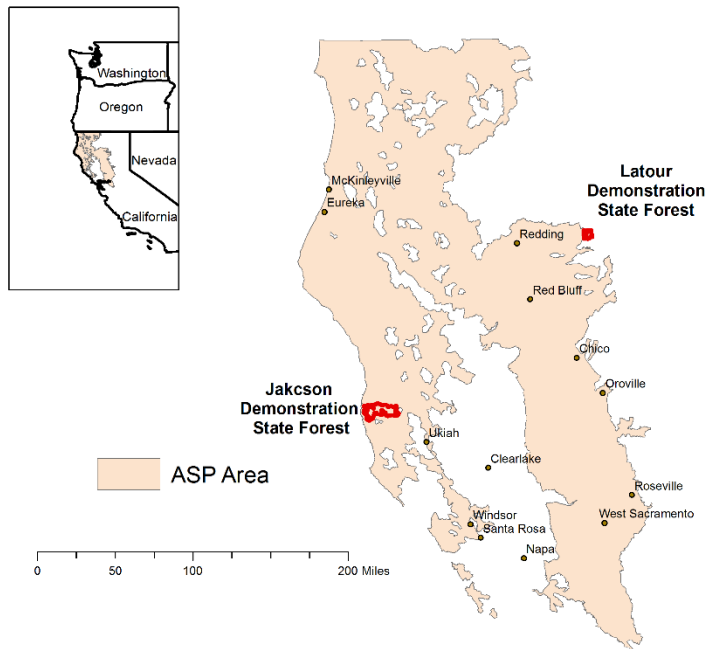


Figure 1: Study area corresponding to the Anadromous Salmonid Protection (ASP). A detailed field campaign will be developed in two Demonstration State Forests (Jackson and Latour).

3. Approach

This investigation will incorporate two scales of inquiry, including a) broad scale analysis of relative controls of physiographic and climatic variables on the perennial flow extent of Class II-L streams as part of the ASP area, and b) a focused field-based analysis of the thermal influence of Class II-L on Class I systems incorporating sites underlain by sedimentary and volcanic geology.

3.1. Broad Scale Analysis

We will conduct a stratified field campaign across the confirmed Anadromous Salmonid Protection (ASP) area (Fig. 1). This area is roughly 39,670 sq-miles and encompasses contrasting geology varying from sandstone, shale, and minor conglomerate in the Coast Ranges to metamorphic and volcanic rocks in the Klamath Mountains, and volcanic rocks in the Cascade Range (Fig. 2). We will select a minimum of 100 Class II-L streams across the main geologic units including as much as possible to represent the range of climatic conditions (Fig. 3). The site selection will take advantage of available LiDAR (Fig. 2) and hydrometric data. For each site we will conduct a physiographic analysis, including calculation of drainage area, catchment slope, and channel profile. In addition, we will identify the topographic channel head according to available techniques (Passalacqua *et al.*, 2010; Clubb *et al.*, 2014) and utilizing available bare ground topography data. At least 75 of the sites will be visited during the summer to assess geomorphic characteristics including average active channel width and channel slope and to map the perennial extent. In addition, we will analyze available hydrometric data to characterize the hydrologic regime in terms of flow duration curves, recession curve analysis, and hydrologic storage (Sayama *et al.*, 2011). This information will be used to formulate a model of channel head initiation and the extent of Class II-L streams across the region. This analysis will

contribute to the formulation of an updated regionalized parameterization of a Class II-L identification system. It is important to note that it is conceivable that individual watercourses that do not meet the Class II-L criteria could have sufficient summertime flow to convey thermal impacts downstream, or support beneficial uses within the Class II watercourse itself and, as such, should be provided Class II-L protection measures. While this is a valid issue, quantification of the spatial extent of such streams across the area of concern is beyond the scope of this study at this time.

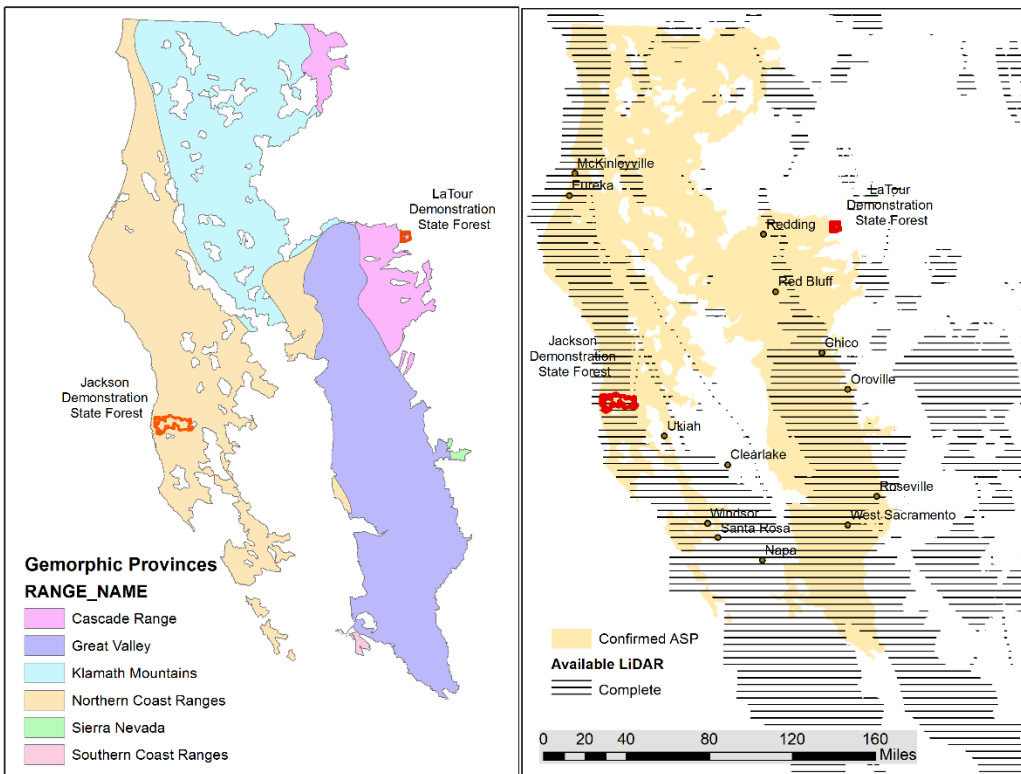


Figure 2: A map of the geomorphic provinces in the study area (left) and available LiDAR (right).

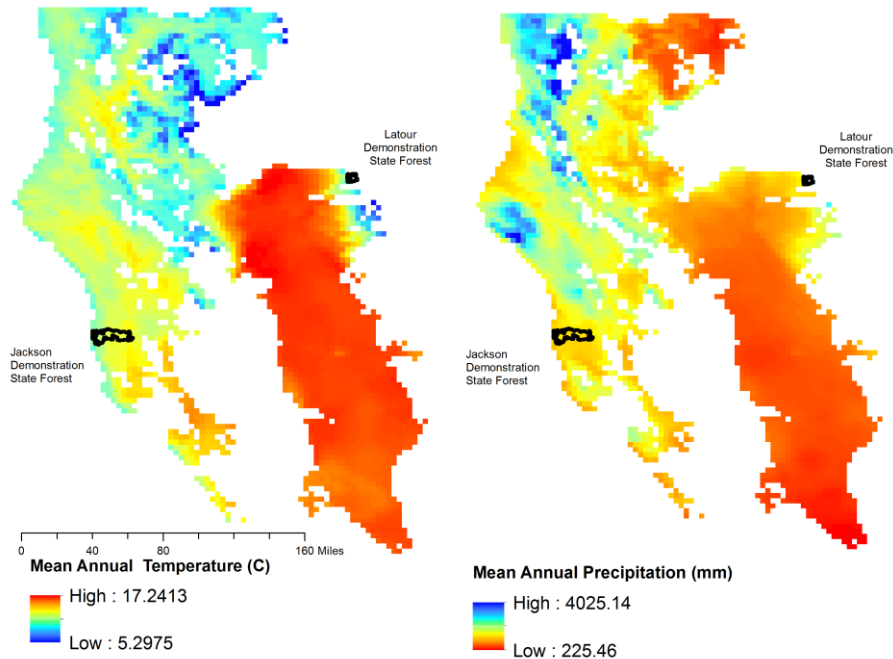


Figure 3: 30year normals (1981-2010) of mean annual tempertaure and precipitation across the study area.

3.2. Longitudinal trends in stream temperature

We propose a pilot study to investigate longitudinal temperature patterns in catchments draining contrasting lithology, with Jackson Demonstration State Forest primarily underlain by sedimentary deposits and LaTour Demonstration State Forest underlain by volcanic rocks. We have tentatively selected sites in Jackson and LaTour Demonstration State Forests (Fig. 2) because they provide readily available access for research, drain contrasting geology, include sites that have not been recently (<15 yrs) harvested, and have available baseline data that will facilitate a robust analysis of trends. We will instrument 6 catchments, including sites in Caspar (Fig. 4) and South Cow Creeks. In each catchment we will deploy Onset TidbiT water temperature data loggers to collect stream temperature data (30-minute intervals). Stream temperature (T_s) loggers will be paired with air temperature (T_a) data loggers to develop direct, local relationships between T_s and T_a . Loggers will be placed approximately 100 m apart along the thalweg of each stream. The sensors will not only provide information about temperature dynamics over time but also have the potential to facilitate determination of the temporal variability of the perennial extent of the network.

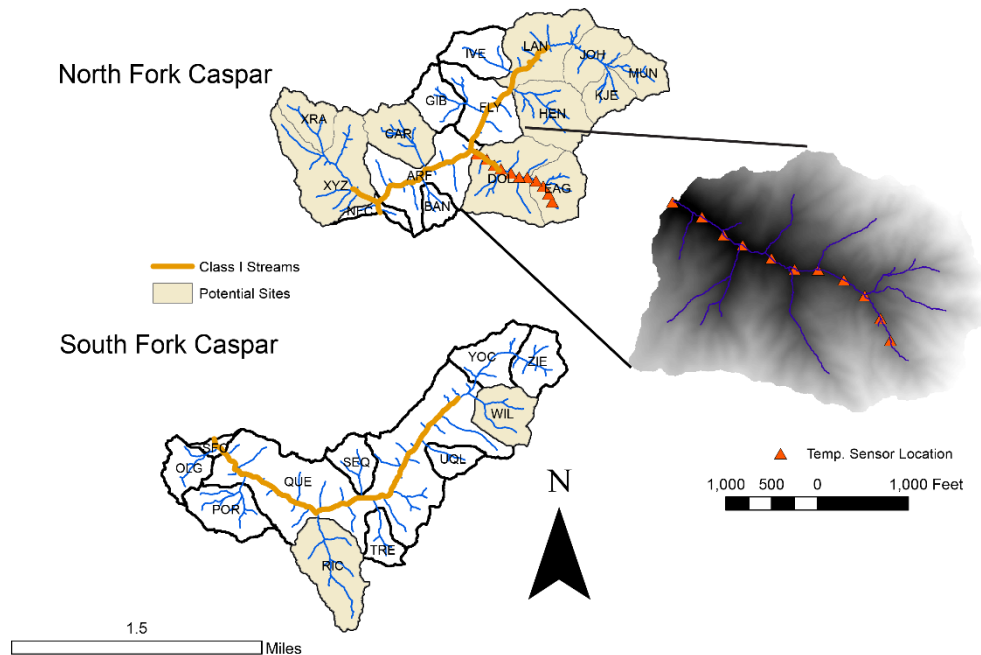


Figure 4. An example of potential sites for thermistors along Caspar Creek in the Jackson Demonstration State Forest.

The comparison between sites in Jackson and LaTour will provide baseline information about heat transmission in the absence of harvesting activities and enable the isolation of the controlling effects of lithology on surface water-ground water interactions. This information will be useful for future studies that target prescription effectiveness, which are beyond the scope of this project but could be part of a second phase.

4. Proposal collaborators

We will collaborate closely with Joseph Wagenbrenner, Research Hydrologist, and Elizabeth Keppeler, Hydrologist, USFS Pacific Southwest Research Station, representatives for the Caspar Creek Experimental Watersheds. We will leverage existing temperature sensors in Caspar Creek to avoid duplicating measurements. We will also collaborate with Drew Coe and Pete Cafferata from CAL FIRE, who will aid in site selection and data acquisition.

5. Timeline

The duration of the project will be 2 ¼ years starting in the summer of 2017 and extending until December 2019 (Table 1). Given that the method for determination of Class II-L is scheduled to sunset on January 1, 2019, preliminary results from the summer 2018 field collection will be delivered in December 2018. The timeline presented in Table 1 indicates the core activities associated with each of the objectives presented in section 2 (a-c).

Table 1: Timeline of the project.

Activity	Academic Year 2017-2018				Academic Year 2018-2019				su19
	su17	fa17	wi18	sp18	su18	fa18	wi19	sp19	
Objective a-b									
Geospatial data compilation									
Hydrometric data compilation									
Site selection									
Field work									
Data analysis									
Objective c									
Site selection									
Instrumentation									
Temperature data collection									
Data analysis									
Thesis defense									
Logger retrieval									

6. Budget justification

- A. **Salaries:** A total of **\$87,319** is allotted for salaries for PI, Segura (\$9,638), co-PI, Bladon (\$9,614), a Masters Student (\$51,939), and a field assistant (\$16,128).
- B. **Fringe benefits:** A total of **\$25,643** was calculated for fringe benefits for all personnel for the duration of the project, and follow approved guidelines. Fringe benefits were calculated at a rate of 44% and 45% for the PI and co-PI, for years 1 and 2, respectively. The MS student fringe benefits are 29%, 31%, and 33% for years 1, 2 and 3, respectively. Fringe for the hourly hired field assistant is 8.31% every year.
- C. **Travel:** A total of \$27,996 is requested for travel for the project duration. This includes:
- Field work 1st year (objective c): \$6,964
 - Field work 2nd year (objectives a and b): \$16,894
 - Travel cost associated to the student participation at the American Geophysical Union Annual Fall Meeting: \$ 1,618
 - Field work 2nd year (objective c): \$2,520
- D. **Materials and Supplies:** A total of \$23,400 is requested for a suite of instrumentation and materials to measure water and air temperature.
- E. **Publication Costs:** A total of \$1,000 is requested to cover the publication cost of one manuscript.
- F. **Computer Services:** A total of \$600 per year is requested to cover the cost of storage of geospatial data.
- G. **Tuition:** The total requested funding for graduate student tuition is \$34,748 for the 2¼ years duration of the project.
- H. **Total direct costs:** Total direct costs of this project are \$201,286
- I. **Indirect costs:** Total indirect costs of this project are \$19,985.
- J. **Total direct and indirect costs:** Total project costs are \$221,271.

7. References

- Beschta RL, Taylor RL. 1988. Stream temperature increases and land-use in a forested Oregon watershed. *Water Resources Bulletin*, **24**: 19-25.
- Bladon KD, Cook NA, Light JT, Segura C. 2016. A catchment-scale assessment of stream temperature response to contemporary forest harvesting in the Oregon Coast Range. *Forest Ecology and Management*, **379**: 153-164. DOI: <http://dx.doi.org/10.1016/j.foreco.2016.08.021>.
- Brown GW. 1969. Predicting Temperatures of Small Streams. *Water Resour. Res.*, **5**: 68-&. DOI: 10.1029/WR005i001p00068.
- Brown GW, Krygier JT. 1970. Effects of Clear-Cutting on Stream Temperature. *Water Resour. Res.*, **6**: 1133-&. DOI: 10.1029/WR006i004p01133.
- Bywater-Reyes S, Segura C, Bladon K. Geology and geomorphology control suspended sediment yield and modulate increases following timber harvest in Oregon headwater streams. *Journal of Hydrology*, **In press**.
- Caissie D. 2006. The thermal regime of rivers: a review. *Freshwater Biology*, **51**: 1389-1406. DOI: 10.1111/j.1365-2427.2006.01597.x.
- Clubb FJ, Mudd SM, Milodowski DT, Hurst MD, Slater LJ. 2014. Objective extraction of channel heads from high-resolution topographic data. *Water Resour. Res.*, **50**: 4283-4304. DOI: 10.1002/2013wr015167.
- Costigan KH, Jaeger KL, Goss CW, Fritz KM, Goebel PC. 2016. Understanding controls on flow permanence in intermittent rivers to aid ecological research: integrating meteorology, geology and land cover. *Ecohydrology*, **9**: 1141-1153. DOI: 10.1002/eco.1712.
- Croke JC, Hairsine PB. 2006. Sediment delivery in managed forests: a review. *Environmental Reviews*, **14**: 59-87. DOI: 10.1139/a05-016.
- Davis LJ, Reiter M, Groom JD. 2015. Modelling temperature change downstream of forest harvest using Newton's law of cooling. *Hydrological Processes*. DOI: 10.1002/hyp.10641.
- Dent L, Vick D, Abraham K, Schoenholtz S, Johnson S. 2008. Summer temperature patterns in headwater streams of the Oregon Coast Range. *Journal of the American Water Resources Association*, **44**: 803-813. DOI: 10.1111/j.1752-1688.2008.00204.x.
- Ebersole JL, Liss WJ, Frissell CA. 2003. Cold water patches in warm streams: Physicochemical characteristics and the influence of shading. *Journal of the American Water Resources Association*, **39**: 355-368. DOI: 10.1111/j.1752-1688.2003.tb04390.x.
- Fritz KM, Johnson BR, Walters DM. 2008. Physical indicators of hydrologic permanence in forested headwater streams. *Journal of the North American Benthological Society*, **27**: 690-704. DOI: 10.1899/07-117.1.
- Fullerton AH, Torgersen CE, Lawler JJ, Faux RN, Steel EA, Beechie TJ, Ebersole JL, Leibowitz SG. 2015. Rethinking the longitudinal stream temperature paradigm: region-wide comparison of thermal infrared imagery reveals unexpected complexity of river temperatures. *Hydrological Processes*, **29**: 4719-4737. DOI: 10.1002/hyp.10506.
- Gomi T, Moore RD, Dhakal AS. 2006. Headwater stream temperature response to clear-cut harvesting with different riparian treatments, coastal British Columbia, Canada. *Water Resour. Res.*, **42**. DOI: 10.1029/2005wr004162.

- Gomi T, Moore RD, Hassan MA. 2005. Suspended sediment dynamics in small forest streams of the Pacific Northwest. *J. Am. Water Resour. Assoc.*, **41**: 877-898.
- Gravelle JA, Link TE. 2007. Influence of timber harvesting on headwater peak stream temperatures in a Northern Idaho watershed. *Forest Science*, **53**: 189-205.
- Gregory SV, Swanson FJ, McKee WA, Cummins KW. 1991. An ecosystem perspective of riparian zones. *BioScience*, **41**: 540-551.
- Grieve SD, Mudd SM, Milodowski DT, Clubb FJ, Furbish DJ. 2016. How does grid-resolution modulate the topographic expression of geomorphic processes? *Earth Surf. Dyn.*, **4**: 627-653. DOI: 10.5194/esurf-4-627-2016.
- Guenther SM, Gomi T, Moore RD. 2014. Stream and bed temperature variability in a coastal headwater catchment: influences of surface-subsurface interactions and partial-retention forest harvesting. *Hydrological Processes*, **28**: 1238-1249. DOI: 10.1002/hyp.9673.
- Hanak E, Lund J, Dinar A, Gray B, Howitt R, Mount J, Moyle P, Thompson B. 2011. *Managing California's Water*. Public Policy Institute of California.
- Jaeger KL, Montgomery DR, Bolton SM. 2007. Channel and perennial flow initiation in headwater streams: Management implications of variability in source-area size. *Environ. Manage.*, **40**: 775-786. DOI: 10.1007/s00267-005-0311-2.
- Keith RM, Bjornn TC, Meehan WR, Hetrick NJ, Brusven MA. 1998. Response of juvenile salmonids to riparian and instream cover modifications in small streams flowing through second-growth forests of southeast Alaska. *Trans. Am. Fish. Soc.*, **127**: 889-907. DOI: 10.1577/1548-8659(1998)127<0889:rojstr>2.0.co;2.
- Kibler KM, Skaugset A, Ganio LM, Huso MM. 2013. Effect of contemporary forest harvesting practices on headwater stream temperatures: Initial response of the Hinkle Creek catchment, Pacific Northwest, USA. *Forest Ecology and Management*, **310**: 680-691. DOI: 10.1016/j.foreco.2013.09.009.
- Leach JA, Moore RD. 2011. Stream temperature dynamics in two hydrogeomorphically distinct reaches. *Hydrological Processes*, **25**: 679-690. DOI: 10.1002/hyp.7854.
- Louen JM. 2016. Hydrologic characteristics of summer stream temperatures in Little Creek and Scotts Creek at the Swanton Pacific Ranch. *Cal Poly San Luis Obispo*, pp: 82.
- Macdonald JS, MacIsaac EA, Herunter HE. 2003. The effect of variable-retention riparian buffer zones on water temperatures in small headwater streams in sub-boreal forest ecosystems of British Columbia. *Canadian Journal of Forest Research*, **33**: 1371-1382.
- MacDonald LH, Coe D. 2007. Influence of Headwater Streams on Downstream Reaches in Forested Areas. *For. Sci.*, **53**: 148-168.
- Madej MA, Currens C, Ozaki V, Yee J, Anderson DG. 2006. Assessing possible thermal rearing restrictions for juvenile coho salmon (*Oncorhynchus kisutch*) through thermal infrared imaging and in-stream monitoring, Redwood Creek, California. *Canadian Journal of Fisheries and Aquatic Sciences*, **63**: 1384-1396. DOI: 10.1139/f06-043.
- McGurk BJ. 1989. Predicting stream temperature after riparian vegetation removal. In: *Proceedings of the California Riparian Systems Conference: protection, management, and restoration for the 1990s*, Abell DL, Technical Coordinator (ed.) Pacific Southwest Forest and Range Experiment Station, Forest Service, U.S. Department of Agriculture, Berkeley, CA, pp: 157-164.
- Montgomery DR, Dietrich WE. 1989. SOURCE AREAS, DRAINAGE DENSITY, AND CHANNEL INITIATION. *Water Resour. Res.*, **25**: 1907-1918. DOI: 10.1029/WR025i008p01907.

- Moore RD, Spittlehouse DL, Story A. 2005. Riparian microclimate and stream temperature response to forest harvesting: a review. *Journal of the American Water Resources Association*, **41**: 813-834. DOI: 10.1111/j.1752-1688.2005.tb04465.x.
- ODEQ. 2004. Oregon's 2004 Water Quality Assessment Section 305(b). Oregon Department of Environmental Quality Water Quality Division, pp: 55.
- Passalacqua P, Do Trung T, Foufoula-Georgiou E, Sapiro G, Dietrich WE. 2010. A geometric framework for channel network extraction from lidar: Nonlinear diffusion and geodesic paths. *Journal of Geophysical Research-Earth Surface*, **115**. DOI: 10.1029/2009jf001254.
- Prosser IP, Abernethy B. 1996. Predicting the topographic limits to a gully network using a digital terrain model and process thresholds. *Water Resour. Res.*, **32**: 2289-2298. DOI: 10.1029/96wr00713.
- Sayama T, McDonnell JJ, Dhakal A, Sullivan K. 2011. How much water can a watershed store? *Hydrological Processes*, **25**: 3899-3908. DOI: 10.1002/hyp.8288.
- Story A, Moore RD, Macdonald JS. 2003. Stream temperatures in two shaded reaches below cutblocks and logging roads: downstream cooling linked to subsurface hydrology. *Canadian Journal of Forest Research*, **33**: 1383-1396. DOI: 10.1139/x03-087.
- Tarboton DG, Bras RL, Rodriguez-Iturbe I. 1991. ON THE EXTRACTION OF CHANNEL NETWORKS FROM DIGITAL ELEVATION DATA. *Hydrological Processes*, **5**: 81-100. DOI: 10.1002/hyp.3360050107.
- Tucker GE, Bras RL. 1998. Hillslope processes, drainage density, and landscape morphology. *Water Resour. Res.*, **34**: 2751-2764. DOI: 10.1029/98wr01474.
- Wang D, Wu L. 2013. Similarity of climate control on base flow and perennial stream density in the Budyko framework. *Hydrology and Earth System Sciences*, **17**: 315-324. DOI: 10.5194/hess-17-315-2013.